

Contribution of Magnetic Resonance Soundings
for characterizing water storage
in the unsaturated zone of karst aquifers

Naomi Mazzilli*, Marie Boucher[†], Konstantinos Chalikakis*, Anatoly
Legchenko[‡], Hervé Jourde[‡] and Cédric Champollion[§]

**UMR 1114 EMMAH (UAPV/INRA), Université d'Avignon et des Pays de Vaucluse, 33
rue Pasteur, 84000 Avignon, France. naomi.mazzilli@univ-avignon.fr,*

konstantinos.chalikakis@univ-avignon.fr

*†UMR 5564 LTHE (IRD/UJF/CNRS/G-INP), BP 53, 38041 Grenoble, France.
marie.boucher@ird.fr, anatoly.legtchenko@ujf-grenoble.fr*

*‡UMR 5569 HSM (CNRS/IRD/UM2), Université Montpellier 2 CC MSE, 34095
Montpellier cedex 5, France. herve.jourde@univ-montp2.fr*

*§UMR 5243 (CNRS/UM2) Géosciences Montpellier, Université Montpellier 2 CC 60,
34095 Montpellier cedex 5, France. cedric.champollion@univ-montp2.fr*

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ABSTRACT

Understanding the role of the unsaturated zone in aquifer recharge and contaminant attenuation processes is a major challenge for the protection and management of karstic water resources. We present the potential of the MRS (Magnetic Resonance Soundings) geophysical method for characterizing the vadose zone of karst aquifers composed of epikarst and infiltration layers. To investigate the hydraulic functioning of the Durzon karst system lo-

cated on the Larzac plateau (Southern France), we used MRS method at sixteen sites. The MRS results have been compared to available geological information and to core water content measurements. The remarkable spatial variability of the MRS response observed in the study area makes it possible to determine ranges of water storage properties in relation to the lithology of the investigated carbonate formations (dolomite, marly and siliceous limestone). All soundings show either constant or increasing MRS water content with depth, which demonstrates that the infiltration zone may be the major water storage entity for permanent water storage, with important consequences for recharge quality and quantity. These results show the feasibility and potential of the MRS method for the characterization of the karst unsaturated zone and for understanding the vertical distribution of water content, which impacts the overall functioning of karst systems.

Keywords: Magnetic Resonance Soundings (MRS), Surface Nuclear Magnetic Resonance (SNMR), karst hydrology, unsaturated / vadose zone, groundwater, geology, carbonate

INTRODUCTION

Groundwater in karst aquifers accounts for approximately 20% of the world's freshwater resources (Ford and Williams, 2007). This vital resource is often highly vulnerable to contaminants that infiltrate through the unsaturated zone of the karst. Understanding karst characteristics and functioning is a prerequisite to efficiently protecting and managing the water resource contained in karst systems.

Karst aquifers are characterized by a highly heterogeneous and organized structure, which results in a specific hydrodynamic behavior. Karst conceptual models generally describe a similar common structure composed of (Perrin, 2003; Sauter et al., 2008; Tritz et al., 2011, e.g.): (i) a soil and epikarst zone, (ii) an infiltration zone, and (iii) a saturated zone (see Figure 1). The relative importance of these compartments in the hydrological functioning depends on rock texture and structure in addition to the genesis of the karst aquifer. The epikarst is the uppermost zone of exposed karstified rocks, where permeability due to fissuring and diffuse karstification is substantially greater than that of the underlying infiltration zone (Klimchouk, 2004). Depending on local conditions, the soil and epikarst zone may be negligible or absent. Thickness of the epikarst depends on lithology and on the geomorphological history of the rock. It is commonly estimated to range from a few meters to 10-15 m (Klimchouk, 2004). The unsaturated zone consists of both the soil - epikarst and infiltration zones. The infiltration zone connects the epikarst to the saturated zone. Therefore, epikarst drainage is controlled by the vertical hydraulic conductivity of the underlying infiltration zone.

The unsaturated zone plays a key role in karst recharge and contaminant attenuation processes (Lastennet et al., 1995). Recent work emphasizes the water storage role of the epikarst in the karst system (Perrin, 2003; Aquilina et al., 2006; Williams, 2008; Hartman et al., 2012). Indeed, low permeability of the infiltration zone combined with an irregular pattern of vertical conduits may result in water storage within the epikarst system. Significant water storage may also occur within the infiltration zone as shown by carbon 13 investigations on the 600m-thick unsaturated zone of the Fontaine de Vaucluse hydrosystem in southern France (Emblanch et al., 2003). Spring water that has had negligible residence time within the saturated zone provides up to 47% of the total discharge of this karst aquifer outlet. The characteristics of the unsaturated zone are thus recognized as an important factor in a karst groundwater vulnerability assessment.

Multicriteria vulnerability assessment approaches produce vulnerability maps based on a combination of different parameters (factors) that are assumed to govern the flow behavior. For karst aquifers, the lithology, thickness and fracturing of the unsaturated zone as well as the capacitive function of the epikarst are recognized as key factors in vulnerability index assessment (Marin et al., 2012) but direct, quantitative assessment of storage in the unsaturated zone remains elusive.

Geophysical methods can provide non-invasive insight into sub-surface properties. Even so, karst systems remain a difficult environment for geophysical exploration. An evaluation of surface-based geophysical methods applied to karst system exploration can be found in Chalikakis et al. (2011). Compared to other non-invasive geophysical methods, the magnetic resonance sounding (MRS) method (Shirov et al., 1991; Legchenko et al., 2004; Behroozmand et al., 2014, e.g.) can directly sense groundwater and consequently, MRS is well suited to hydrogeological characterization. There is a strong theoretical basis for the relationship

between MRS parameters and hydrological parameters (total and effective porosity, specific yield, hydraulic conductivity) (Shirov et al., 1991; Lubczynski and Roy, 2005; Legchenko et al., 2004, e.g.). The broad investigation scale of MRS (several cubic decimeters to cubic hectometers depending on the set-up (Legchenko et al., 2004)) is also relevant for hydrodynamic investigations (Lachassagne et al., 2005). The MRS method was initially developed for characterizing the saturated zone. Most MRS applications on karst systems (Vouillamoz et al., 2003; Boucher et al., 2006b; Girard et al., 2007; Legchenko et al., 2008a; Pérez-Bielsa et al., 2012, e.g.) have thus been devoted to the investigation of saturated zone properties. However, water can be detected in the unsaturated zone in weakly magnetic environments: e.g. in carbonate rocks or in sandy formations (Legchenko et al., 2002; Miché et al., 2003; Lubczynski and Roy, 2005; Boucher et al., 2006a; Vouillamoz et al., 2012; Walsh et al., 2014; Costabel and Gunther, 2014).

This paper is an investigation of MRS potential for the characterization of the unsaturated zone of karst hydrosystems. The main questions addressed are:

- does water storage within the unsaturated zone of karst yield quantifiable MRS measurements ? If so,
- is MRS suitable for quantifying the spatial variability of unsaturated zone water content ?
- can MRS provide additional insight into karst structure and functioning ?

For that purpose, MRS soundings were used to investigate 16 experimental sites in the Larzac plateau area. The proposed methodology is based on both geological and hydrogeological analyses of the experimental sites. To the author’s knowledge, this study constitutes

the first application of the MRS method to the unsaturated zone of karst aquifers.

STUDY AREA

MRS investigations were performed in the Larzac karstic plateau, Southern France (see location in Figure 2). The choice of this area was motivated by the wide diversity of unsaturated zone characteristics in the study area.

Environmental setting

Primary climate influences in the study area are the Mediterranean Sea, the Atlantic Ocean, and altitude effects. The altitude of the plateau ranges from 560 to 920 m a.s.l. Mean annual rainfall in the area is approximately 1000 mm based on the 1990-2010 period, with the highest interannual average rainfall occurring in the south-western part of the plateau. Maximum rainfall occurs in autumn and winter. Historical changes in human activities, principally dwindling sheep grazing, cultivation shift and woodland exploitation have been postulated as having major impacts on vegetation dynamics over the past decades (Lepart and Debussche, 1992; Kunstler et al., 2007). At present, urbanization is minor and most economic activities are associated with sheep and cow farming, and agriculture. The landscape consists of a mosaic of croplands, open and encroached grasslands and pine and oak woods.

Hydrogeological setting

The Larzac plateau is located within Lias to Malm Jurassic limestones and dolomites, 200- and 400-m thick respectively, separated by a ~ 100 -m-thick impermeable formation of Upper Lias marls (see Figure 2). Regional structure is nearly horizontal, cut by major E-W and

ENE-WSW faults continuously active since the Lower Jurassic (Paloc, 1972; Plagnes and Bakalowicz, 2001). A detailed description of the stratigraphy of the Dogger and Malm series can be found in e.g. Charcosset (1998) and Bruxelles (2001). Springs discharge at the base of the carbonate series. The main outlet is the Durzon spring (533 m ASL), which has average outflow of 1.4 m³/s based on the 2002 - 2008 period.

Superficial deposits are primarily represented by clays with occasional siliceous deposits stemming from the alteration of early Bajocian limestones, in particular due to ghost rock weathering processes (Bruxelles, 2001). This alteration product is transported and collected into depressions that have insufficient vertical drainage. Dolomitic sand derived from the alteration of dolomite is present over most dolomite outcrops that lack clay cover. Superficial deposits (composed of clays and dolomitic sand) tend to clog the vertical fractures and karstic openings and therefore limit the efficiency of vertical drainage. Clay cover resulting from alteration processes (terra rossa) tends to promote runoff and therefore focuses concentrated infiltration towards the limestones.

The primary processes involved in the genesis of the Larzac karst include: (i) in-situ alteration, (ii) crypto-karstification (i.e., karstification under cover), (iii) karstification due to the presence of high hydraulic gradients associated with the incision of the Grands Causses canyons (Bruxelles, 2001). Retrogressive and recent karstification associated with most outflows has established a high karstification gradient between the upstream and downstream parts of main springs catchments. At the center of the study area, limited vertical drainage may cause saturation of the epikarst during intense and lengthy rainfall periods, thus leading to the formation of temporary lakes and rivers at the surface (Bruxelles and Caubel, 1996). The chemical signature of the Durzon spring water indicates long residence time within the unsaturated zone (Bondu-Crozol, 2012). According to observations based on time-lapse

microgravimetry surveys (Jacob et al., 2009, 2008), seasonal water storage variations in the unsaturated zone may reach 550 mm.

MRS sites

A total of 16 sites were selected for MRS investigations, based on the following two criteria: (i) diversity and spatial representativeness of the hydrogeological and geomorphological setting, and (ii) ambient electromagnetic noise conditions, which can affect MRS data quality (Plata and Rubio, 2002; Girard et al., 2005; Legchenko, 2007; Chalikakis et al., 2008). All selected sites are located within, or close to, the Durzon spring catchment. Average thickness of the unsaturated zone in the study area is ~ 100 m.

Ambient electromagnetic noise conditions were assessed using preliminary noise mapping using noise tester from IRIS Instruments (Bernard, 2007). On the Larzac plateau, the major anthropic source of electromagnetic noise is a high voltage power line, which causes significant disturbances to the MRS signal up to a distance of 1 to 2 km from the line. Low voltage power lines and electric fences set up for livestock farming also disturb the MRS signal to a distance of a few hundred of meters from the noise source. Raw noise measurement were used to select sites with both stable and low raw noise conditions (< 120 nV for a 6×6 m 10 turns loop at the 2000 Hz frequency). At each sounding site, we then used the raw noise measurer to check the magnitude and variability of ambient noise at the site and thus select the type of loop (either square or eight square) and the orientation of the eight square loop (if selected) that would minimize disruptions due to surrounding power lines (Trushkin et al., 1995).

We assessed the hydrogeological and geomorphological setting of the experimental sites

based on the approach proposed by the European COST Action 620 on “Vulnerability mapping for the protection of carbonate (karst) aquifers” (Daly et al., 2002). The following factors were retained for site characterization: (i) lithology, (ii) fracturing and karstification, which may favor infiltration flow, (iii) vegetation, slope and clay cover, which may favor runoff. Note that karst infill material may clog fissures and therefore reduce the infiltration capacity of the formation. However, a qualitative and quantitative characterization of karst infill material is limited to sites with possible access to the subsurface (either through drillings or cave exploration). In this study, we consider that: (i) ruiniform dolomite may yield sand infill material, and (ii) clay cover may yield clay infill material.

The site characteristics are summarized in Table 1. Most MRS soundings investigate the ruiniform dolomite formation that outcrops throughout the study area. Sites H2, H3, and Menu were selected to monitor the temporal variability of the MRS signal because of low noise conditions and relatively high water storage variations expected at these sites. Gravity variations of $15.8 \mu\text{gal}$, corresponding to a variation of an equivalent water slab thickness of 37 cm, have been monitored at the center of the H3 site (Jacob et al., 2008). The H2 site is located at a distance of 100 m from H3, in a sinkhole. The Menu site is located upstream of the Durzon spring catchment, close to the temporary lake area.

MATERIAL AND METHOD

Principle of Magnetic Resonance Soundings

Basics

The Magnetic Resonance Soundings method is based on the resonance behavior of the magnetic moments of protons in the groundwater molecule. These moments are able to absorb and to emit energy from an electromagnetic field at a specific frequency (Larmor frequency) that ensures selective sensitivity of the method to groundwater. The electromagnetic field is produced by a surface loop energized by a pulse of oscillating current. After the current pulse is terminated, the energy absorbed by magnetic moments is emitted back. This response can be measured by a receiving loop on the surface. The MRS signal parameters (initial amplitude e and T_2^* relaxation time) are derived from the MRS signal envelope assuming a mono-exponential decay. One sounding is composed of ten to twenty MRS signals measured for different values of the pulse moment. Ambient electromagnetic noise is recorded over a few hundreds of milliseconds before the current pulse is transmitted. It makes it possible to estimate the signal to noise ratio (S/N) as a ratio of the amplitude of the magnetic resonance signal to a mean of the electromagnetic noise. To improve the signal to noise ratio, measurements are repeated and stacked (usually tens to several hundred times, depending on the ambient noise magnitude). The basics of the method are described in details in e.g. Legchenko and Valla (2002).

Hydrogeological interpretation of MRS results

The parameters derived from the magnetic resonance signal are the volumetric MRS water content (θ_{MRS}), the longitudinal relaxation time (T_1) and the observed transverse relaxation time (T_2^*) versus depth. Relaxation times T_1 and T_2^* depend on the mean distance between the water molecule and the internal rock surface: the shorter the distance, the shorter the relaxation times (Legchenko et al., 2004). T_1 and T_2^* are thus affected by pore size (Grunewald and Knight, 2011) and by saturation because in unsaturated media water remains near the rock surface due to capillary forces (Lubczynski and Roy, 2005). The mean distance of water molecules to pore surface matters only if the pore space dimension is smaller than a few millimeters. In larger pores or conduits MRS measures the relaxation time of bulk water.

Measuring T_2^* is easier and faster than measuring T_1 because the T_1 assessment requires the use of a two-pulses protocol with a variable delay between pulses (Dunn et al., 2002; Legchenko et al., 2004). However, T_2^* is also affected by local heterogeneities of the local magnetic field and hence is often a less reliable parameter than T_1 (Farrar and Becker, 1971). In the Larzac area, where the subsurface is composed of limestone and dolomite that have low magnetic susceptibility, T_2^* relaxation time can be safely used. In carbonate rocks, typical values of the relaxation time T_2^* range from 80 to 130 ms in the absence of large or multiple karst voids (Legchenko et al., 2002; Miché et al., 2003; Boucher et al., 2006b). Otherwise, it is generally higher than 400 ms (Legchenko et al., 2002; Vouillamoz et al., 2003).

Due to instrumental dead time of the applied equipment, MRS cannot detect a certain portion of the water (Boucher et al., 2011). MRS water content θ_{MRS} is defined as the volume of detected water per unit of sampled volume. MRS water content is thus lower

than or equal to the total volumetric water content of the investigated rock (Shirov et al., 1991). The fraction of undetectable water is characterized by short relaxation time (≤ 40 ms). Relaxation times lower than 40 ms are generally associated with bound water and sometimes capillary water (Vouillamoz et al., 2012). The threshold of undetectable water for MRS depends on the magnetic properties of the rock. Vouillamoz et al. (2012) proposes an MRS apparent cutoff time approach for differentiating gravitational water and capillary and bound water. In this study, MRS results were compared to the water content measured in core samples.

Effect of electrical conductivity on the MRS signal

The electrical conductivity of the surrounding media can affect the MRS response (Trushkin et al., 1995; Legchenko et al., 2008b, e.g.). However, only layers with resistivity lower than 10 $\Omega\cdot\text{m}$ have a significant impact on the MRS signal (Valla and Legchenko, 2002) and should be considered in the inversion process. Layers with resistivity higher than 100 $\Omega\cdot\text{m}$ have a negligible effect on the MRS signal. Between 100 and 10 $\Omega\cdot\text{m}$, the signal can be slightly affected, particularly if the conductive layer is thick. In carbonate rocks, resistivity is generally high. Previous studies on the Larzac plateau (Valois, 2011) have shown that resistivity is generally higher than 100 $\Omega\cdot\text{m}$ except in clay-covered areas (resistivity may reach 30 to 50 $\Omega\cdot\text{m}$ in these clays). At MRS sites with clay cover, electromagnetic mapping (EM-34) with a 20-m coil spacing was used to evaluate the homogeneity of the electrical conductivity distribution. For most soundings, apparent electrical resistivity measured in both horizontal and vertical dipole configurations was above 100 $\Omega\cdot\text{m}$. At Menu and Fig sites, apparent electrical resistivity was below 100 $\Omega\cdot\text{m}$ at some stations. Electrical resistivity tomography performed at the Menu site by Valois (2011) shows a thin layer of average 50

$\Omega\cdot\text{m}$ resistivity from 2 to 5 m below the surface. Electrical resistivity tomography performed at the Fig site (this study) shows a layer of average 40 $\Omega\cdot\text{m}$ resistivity to a depth of 15 m below the surface. MRS inversion on Menu and Fig sites was performed two ways: including and not including the resistivity model derived from ERT measurements. In retrospect, the influence of resistivity on the MRS signal was negligible in this area.

Field setup

MRS data acquisition have been conducted over six field seasons: July 2009, April 2010, October 2010, May 2011, August 2011 and January 2012. During the first three MRS exploration campaigns from 2009 through 2010, NUMIS^{LITE} equipment from IRIS Instruments was used. NUMIS^{PLUS} systems have been used subsequently. The signal record duration was set to the usual value of 240 ms. The coincident transmitting / receiving loop configuration which is the most suitable for 1-D application was implemented (Valla and Legchenko, 2002). At most sites, eight square loops (composed of two squares of 40 m side) was required because of the high noise level. By using this loop layout, the orientation of the eight was chosen so as to minimize disruptions due to surrounding power lines. When the ambient noise was low, a 80×80 m² square loop was selected to increase the investigation depth.

MRS signal processing

Prior to acquisition, we used the noise in time domain to detect peaks and we checked the noise frequency spectra for harmonic noise. Spikes were not detected at any of our monitoring sites. Most noise came from harmonics. The post-processing procedure consisted of: (i) stacking, (ii) 14-Hz band pass filtering, (iii) 50-Hz notch filtering, (iv) exponential fitting

of the MRS signal, (v) estimation of MRS parameter uncertainty. The number of stacks was set between 100 and 400 depending on ambient noise. As a result, sounding duration ranged from 4 to 8 hours, which limited soundings to only one or two per day. Estimated maximum resolution depths ranged from 32 to 59 m depending on the measuring setup and the signal to noise ratio (for further explanation of scripting in maximum resolution depths estimation, see p.16). Characteristics of the soundings are summarized in Table 2.

MRS signal uncertainty

MRS signal measurements are sensitive to ambient electromagnetic noise (either natural or anthropic). The noise can be decreased through the application of signal stacking during field measurements, and also by applying numerical filtering techniques (Legchenko, 2007). During the sounding, the current noise amplitude can be estimated from extra noise readings recorded before each energizing.

In this study, we considered MRS signal parameter sets (i.e., e and T_2^*) to be acceptable if they yielded an increase in the root mean square error (RMSE) of less than 10% compared to the optimum fit. In this study, we considered as acceptable all MRS signal parameter sets (signal amplitude e and T_2^* relaxation time) that yielded a 10% increase in the root mean squared error (RMSE) fitting error as compared to the optimal RMSE. Uncertainty bounds are derived from the min and max values of MRS parameters (e and T_2^*) that provide an acceptable fit (up to -10% from the optimal RMSE). The uncertainty estimation procedure is illustrated in Figure 3. The 10% threshold was set as follows. For each pulse, we generated a set of synthetic MRS records composed of measured MRS signal (e , T_2^*) with added random noise of the amplitude corresponding to noise measurements before the pulse. This synthetic

signal was processed using the same filtering as for the real data. The standard deviation of the estimated e and T_2^* represent uncertainty in the signal parameters. We have found that for presented data set the standard deviation was varying between 6 and 9% of the signal amplitude and we prescribe the safe value of 10% as the uncertainty estimate.

Inversion approach and model uncertainty

Least-squares inversion of the MRS signal was performed using SAMOVAR software 11×62 based on amplitude processing of complex signals using the Thikhonov regularization (Legchenko et al., 2004). In the absence of magnetic heterogeneities the T_2^* estimate is considered reliable, so we used T_2^* in our survey (Legchenko et al., 2002). The goal of the inversion process is to find the $(\theta_{\text{MRS}}, T_2^*)$ distribution with depth (i.e., the 1-D model) that best accounts for the recorded MRS signal. The workflow inversion procedure used in this study can be summarized as follows (Legchenko and Valla, 2002): (i) the inversion of the initial amplitude leads to the water content distribution $\theta_{\text{MRS}}(t, z)$, (ii) each $\theta_{\text{MRS}}(t, z)$ is fit by an exponential function, thus providing the $T_2^*(z)$ values.

The solution of the inversion problem is not unique because: (i) the measured data is corrupted by noise, and (ii) buried water layers that have equal water volumes (product of water content per layer thickness) can yield equivalent MRS responses (Legchenko and Valla, 2002). The inversion problem must therefore be constrained either by reducing the number of model parameters (block inversion, i.e. small number of discrete layers having individual water contents and boundary depths), or by constraining the model parameters (smooth inversion, i.e. larger number of layers with fixed boundaries but minimized variation of water content between adjacent blocks) (Yaramanci and Hertrich, 2007). Note that

inversion can also be constrained by independent data (stemming from either geological or other geophysical techniques), when pertinent and available. We targeted the soundings performed in this study to different geological units and we had no information regarding lithological variations with depth. We therefore use a smooth, singular value decomposition (SVD)-based inversion procedure (Muller-Petke and Yaramanci, 2008; Legchenko, 2013) that requires no a priori knowledge on the data :

- (i) model layers (number and bounds) are selected with respect to singular values distribution (Legchenko and Shushakov, 1998). Assuming a zero smoothing factor, the number of layers is set so that without regularization the resolution matrix is close to the identical matrix.
- (ii) then the smoothing factor is chosen as a compromise between the smoothness of solution and the accuracy of data fit. In general, a larger smoothing factor makes it possible to stabilize the inversion but degrades resolution and datafit.

The SVD inversion also provides an estimate of the solution uncertainty given by the standard deviation of the inverse model. To estimate the vertical resolution of the inversion, we use the primary diagonal of the model resolution matrix (“depth resolution”) which shows how well the particular layers can be resolved. Values close to one correspond to well resolved layers in the inversion results. We also estimate the maximum resolution depth, which is arbitrary defined as the maximum depth of layers with a depth resolution of 0.5. The maximum resolution depth is always smaller than the sounding investigation depth.

RESULTS AND DISCUSSION

Water content measurability in the karst unsaturated zone using MRS

The soundings results are summarized in Table 2. MRS data quality is estimated based on the average S/N (taking into account both external and instrumental noise) and the average external noise to instrumental noise ratio (EN/IN) (Legchenko, 2007). The S/N and EN/IN estimates are both derived from stacked and filtered measurements. For the EN/IN estimation, the instrumental noise for the NUMIS instrument after stacking and filtering is assumed to be 5 nV (Legchenko, 2007). The higher the S/N, the higher the reliability of the MRS survey. Using the selected setup (measuring device and data processing), we considered MRS soundings to be of acceptable quality when either the S/N is ≥ 2 or when the S/N is ≤ 2 and the EN/IN is ~ 1 . For S/N ~ 1 , only the maximum water content can be derived from the MRS signal.

In this study, soundings performed at 14 of the 16 selected sites meet the quality criteria, thirteen of them have an average S/N ≥ 2 , which confirms (i) the applicability of the MRS method to the study area, and (ii) the fact that water storage in the unsaturated zone of karst may yield measurable MRS signals.

Temporal variability of MRS measurements

MRS measurements have been repeated at three sites (Menu, H2 and H3) since the beginning of the project (2009). Figure 4 shows non-inverted MRS parameters as a function of the energizing pulse for these sites. No significant temporal change in the signal amplitude and relaxation time is observed, which means that water storage variations are below the MRS detection threshold. We define MRS water storage over the thickness Δz as the product of

$\theta_{MRS} \times \Delta z$. Sensitivity analyses of the inverted MRS signal show that MRS water storage variations within the upper 40 m are less than ± 220 mm at the Menu site, less than ± 250 mm at the H2 site, and less than ± 170 mm at the H3 site.

These results are compared to catchment-scale water storage changes between the MRS fieldwork seasons. Catchment-scale water storage is estimated based on simulation results from the conceptual rainfall-runoff model for the Durzon catchment proposed by Tritz et al. (2011). Water storage changes as a function of time are reported in Figure 5. It is seen that the catchment-scale global water storage changes between the MRS campaigns are less than 130 mm equivalent water thickness, which is roughly equal to the MRS detection threshold. Based on these results, we consider the temporal variability of the MRS signal to be minor in the intervals between field campaigns, therefore, all sites can be compared to each other, independently of the date when the data was collected.

It has been shown (Descloitres et al., 2008) that MRS can be used in time-lapse mode for investigating temporal variations of water content. However, to use MRS, the investigated rocks being studied should have a high specific yield and constitute a thick layer with variable water content. In limestones, specific yield is usually low for the matrix but may be high for large pores and karstic structures. Temporal monitoring of water storage variations in karst using MRS should be feasible provided that both extremely low and extremely high water levels are targeted. Karst aquifers are usually highly reactive systems so that field work targeting high water levels should be conducted quickly after rainfall events. To target high water levels, the maximum delay between rainfall events and field work depends on the site. In the case of the Durzon system, a peak flood lasts ~ 7 days (Figure 5a). We could not conduct MRS measurements during these periods, thus no water content variation was observed in this study.

Hydrogeological interpretation and consistency of MRS water content

In this study, the hydrogeological interpretation of MRS water content is supported by the comparison between MRS water content and the analysis of core samples at the H3 site.

Laboratory measurements were done to determine the water content of core samples from one of the drillholes installed at the center of the H3 site. To minimize water loss by evaporation, core samples were sealed immediately after collection. The samples were weighed, dried at 80°C for 48 h, and weighed again to calculate the initial water content. Sample volume was derived from the volume of fluid displaced immediately after sample immersion.

MRS signal and inversion results for the H3 site are shown in Figure 6. The results of core sample analysis are presented in Figure 7. Visual observations of the core show high macro-porosity to a depth of 5 m. Competent dolomite with fine macro-porosity is found from 5 to 14 m. Below 14 m, the grade of dolomite alteration varies widely and macro-porosity appears to diminish with depth (see simplified log in Figure 7a). Water content ranges from 3 to 5.5 % for most 0 to 20 m core samples, higher values are associated with shallow samples. Below 20 m, core water content ranges from 6 to 12.5 %. MRS model uncertainty was estimated using the SVD method. The favorable correspondence between measured and modeled amplitudes (Figure 6) confirms adequate quality of the MRS inversion. MRS model and core sample water contents are depicted in Figure 7b, modeled T_2^* is shown in Figure 7c. As discussed above, MRS equipment does not register a portion of the signal due to the short relaxation time of capillary water. Figure 7b shows that the ratio of MRS and core water content is not constant with depth, which means that the amount of invisible water varies along the lithological profile. This observation can be interpreted as follows:

- Above 10 m the MRS water content underestimates the sample results, which may be due to either increased dolomite weathering and thus higher capillary or bound water content or to decreased resolution and small MRS signals coming along with small MRS pulses in the first 10 m.
- Between 10 and 20 m, MRS and core water content are in good agreement. Here the amount of invisible water is nearly negligible. Maximum MRS resolution is obtained for this depth range, as shown in Figure 7d.
- Below 20 m, MRS slightly underestimates the mean water content measured in the samples, which show a high degree of heterogeneity in water content. Possible reasons for these underestimations include 3-D effects or an increased uncertainty of MRS results due to decreased resolution (Figure 7d). Nevertheless, MRS shows the correct trend of increasing water content with depth.

Analysis of the spatial variability of the unsaturated zone water content

Significant variability is observed in both the initial amplitude and the relaxation time of the MRS signals. For example, for eight square loops measured initial amplitudes range from less than 20nV up to 220 nV whereas average noise is about 15 nV. MRS parameters make it possible to discriminate between different geological settings. For all loop settings, average T_2^* is higher than 120 ms for soundings performed in the ruiniform dolomite environment whereas it is lower than 100 ms for siliceous limestones. It was impossible to estimate T_2^* of marly limestone because of the small signal amplitude (below 20 nV).

MRS inversion results are summarized in Table 3 and illustrated in Figure 8 for selected soundings. No exponential decay could be fit to marly limestone records, and we provide only

the maximum water content based on a forward fit to the measured noise. Figure 9 provides a summarizing view of the inversions. It shows the relaxation time $T_2^*(e_{max})$ associated with maximum signal amplitude, as a function of the water content θ_{mean}^{0-25m} . We focused on the upper part of the soundings (depth < 25 m and pulses < 2500 A.ms) because of the best resolution and these depths/pulses values are reached by all soundings. The relaxation time $T_2^*(e_{max})$ associated with maximum signal amplitude characterizes the lithological properties of the rock (including possible alteration and clay infilling) since the relaxation measured in a specific pore space always decreases with its saturation degree as shown in recent laboratory (e.g. Ioannidis et al. (2006); Costabel and Yaramanci (2011); Mohnke et al. (2015); Boucher et al. (2011)) and field studies (e.g. Walsh et al. (2014); Costabel and Gunther (2014)). On the other hand, θ_{mean}^{0-25m} corresponds to the MRS water content. Soundings performed in marly limestones, siliceous limestones and ruiniform dolomite cluster in three different groups, which confirms the lithological control of the MRS water storage. The variability observed in both $T_2^*(e_{max})$ and θ_{mean}^{0-25m} also indicates that lithology is not the only factor involved in MRS water storage.

Further investigation of the relationship between MRS water storage and flow factors such as those used in vulnerability approaches is hindered by: (i) the number of soundings performed, which does not permit statistical analyses such as principal component analysis and (ii) the uncertainty regarding MRS results, which is in part due to equivalency issues in the inversion of the MRS signal. The fact that flow factors that are traditionnally used in vulnerability assessment methods (e.g. fracturation or karstification features, slope, clay cover) could not be used to predict the relative water content of the investigated sites illustrates (i) the complexity of the processes involved, and (ii) the specific contribution of MRS. Future work focused on this relationship should be performed in low noise environments

while varying one factor at a time.

Insights into the vertical structure of the karst unsaturated zone

The vertical distribution of water content observed in core samples at one site and generalized as a result of MRS prospecting enable us to improve conceptual models of water storage in karst systems. Indeed, the epikarst is usually assumed to have a large capacitive function as compared to that of the underlying infiltration zone and thus to play a major role in seasonal water storage (Klimchouk, 2004; Williams, 2008). In the Durzon karst system, the dominant role of the epikarst in seasonal water storage is shown by surface-to-depth gravity surveys (Jacob et al., 2009; Deville et al., 2011). Even so, MRS investigations for the most part show constant or increasing water content and relaxation time with depth (Table 3). Assuming a 10 m maximum epikarst thickness (Williams, 2008), it can also be seen that for most soundings, the average MRS water content is higher in the infiltration zone than in the epikarst zone (Table 3). These results suggest that, at the watershed scale, the infiltration zone is the major water storage entity for permanent water storage. This has important consequences for the quantity and quality of recharged water because a high water content in the soil and rock facilitates the piston flow effect (Lange et al., 2010). In the Durzon spring catchment, surface runoff is limited due to the mild relief but the high water content of the infiltration zone may as a result speed up the flood dynamics. The outflow of highly mineralized water during flood events (Bondu-Crozel, 2012) is also consistent with long residence time in the infiltration zone. The infiltration zone's key role in water storage was already shown in the 600-m thick infiltration zone of the Fontaine de Vaucluse karst system (Emblanch et al., 2003). Our findings show that the infiltration zone may also play a key role in karst systems with moderate-thickness (~ 100 m) infiltration zones. This result

cannot be generalized to all karst systems because structure and therefore functioning of karst aquifers varies widely. However, the dominant role of the epikarst, which is often assumed in conceptual models should be tempered.

CONCLUSIONS

This study provides both a methodological contribution to the use of MRS and a contribution to the understanding of water storage processes in the karst unsaturated zone. The key points of the methodological contribution are as follows:

1. The water content within the unsaturated zone of karst does yield quantifiable MRS signals. Furthermore, in the ruiniform dolomite that was studied, undetected water is negligible compared to the total water content. As a result, MRS water content yields a good estimate of total water content.
2. Temporal variation of water content between field seasons is below the detection threshold of MRS because of the rapid temporal dynamics of the catchment, which is typical of karst systems.

Regarding water storage in the unsaturated zone of the Durzon karst aquifer, our study yields the following findings:

1. MRS water content was quantified in different geological settings: resolved MRS water content is below 0.5% in marly limestones, and ranges from 2.5 to 5 % in siliceous limestones and from 2 to 20 % in ruiniform dolomite.
2. In most soundings, MRS water content is either constant or increases with depth, which demonstrates that the infiltration zone may be the major water storage entity

for permanent water storage. This result has important consequences for recharge quality and quantity. In particular, high water content in the infiltration zone may facilitate the piston flow effect and therefore accelerate the flood dynamics.

3. No direct relationship could be observed between MRS water content and factors used in vulnerability assessment studies. Further work investigating this relationship should be performed in low noise environments while varying one factor at a time.

These results show the feasibility and potential of the MRS method for the characterization of the karst unsaturated zone and for the understanding of the vertical distribution of water content, which impacts the overall functioning of karst. In the study of the Durzon system, the MRS results improved the representativeness of local information that was obtained by analysis of core samples and highlighted the major role of infiltration in permanent water storage.

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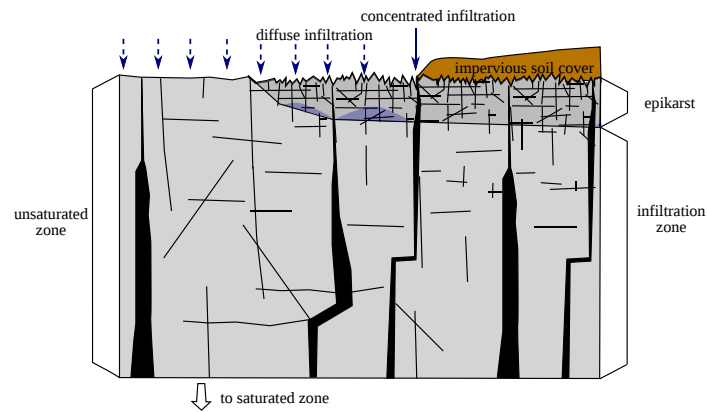


Figure 1: Conceptual model of the unsaturated zone structure and hydrodynamic functioning of karst aquifers, modified after Klimchouk (2004). Note that the sharpness of the transition between epikarst and infiltration zones may vary depending in particular from the homogeneity of the lithographic sequence.

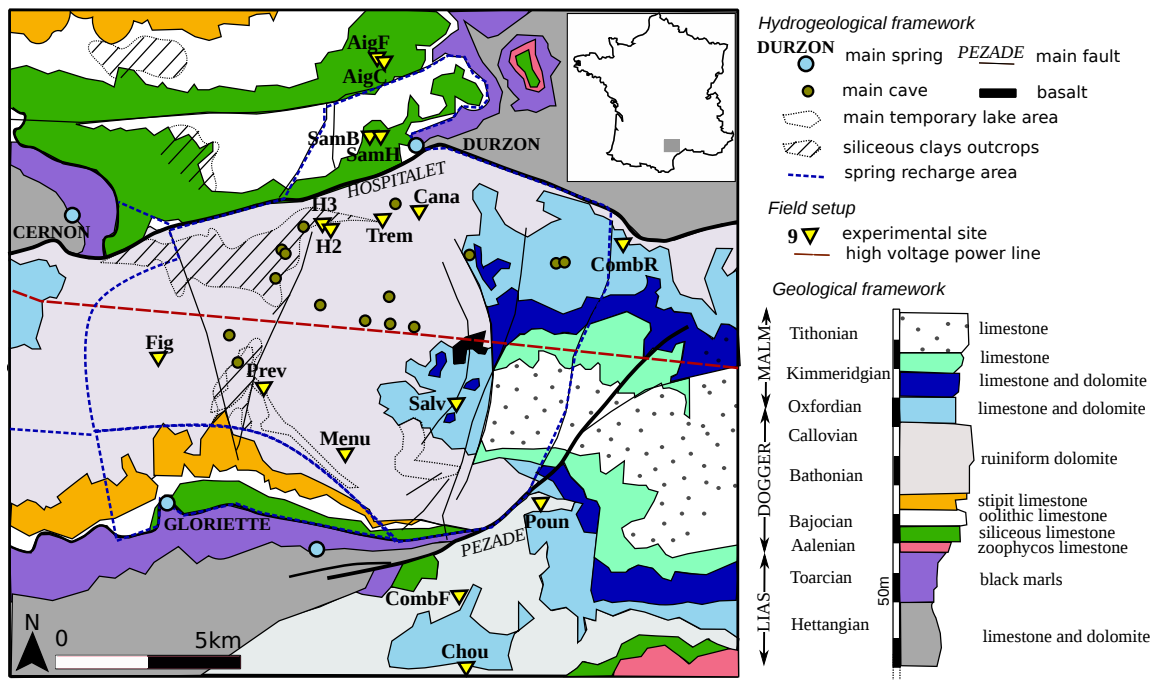


Figure 2: Situation map for the MRS soundings. Geological setting adapted from Bruxelles (2001).

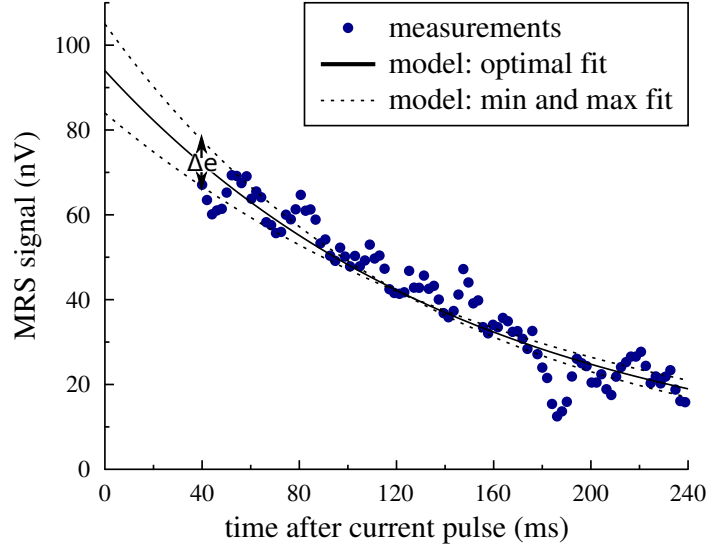


Figure 3: Recorded signal amplitude as a function of time after the current pulse is terminated (Menu sounding, April 2010 campaign, pulse moment equal to 871 A.ms). Note that MRS measurements can only start after an instrumental dead-time of 40 ms which is needed for switching the instrumentation from transmitter to receiver. Blue dots: MRS signal. Solid line: optimal fit of the exponential model. Dashed line: fits yielding a 10% increase from the optimal RMSE. Δe : uncertainty bounds for the signal amplitude.

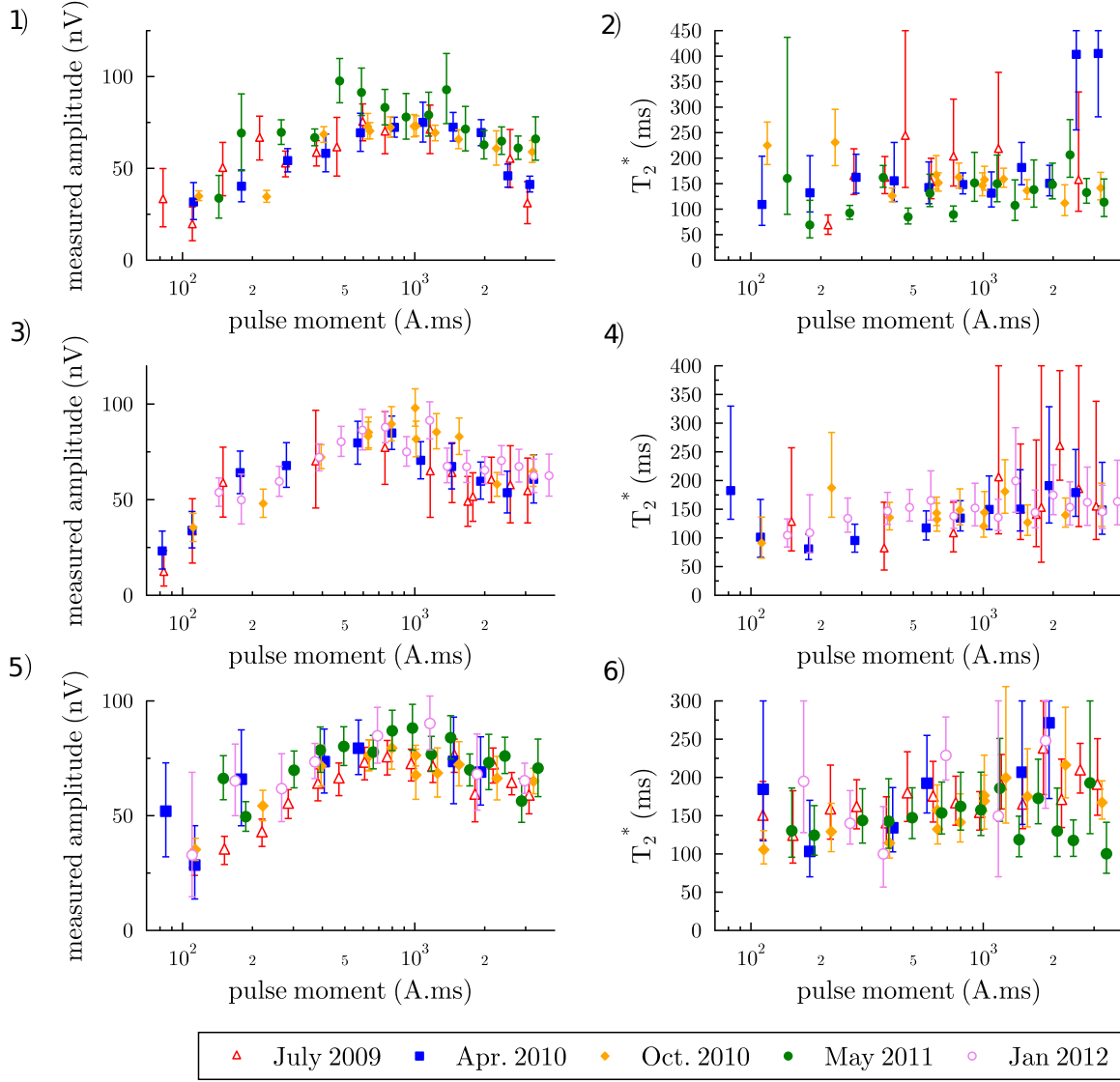


Figure 4: Temporal variability of MRS signals. Signal amplitude e (left) and apparent transverse relaxation time T_2^* (right) of the MRS signals as a function of the energizing pulse, for sounding performed at Menu (graphs 1, 2), H2 (graphs 3, 4) and H3 (graphs 5, 6) sites. Only pulses with a S/N greater than 1.5 have been used for the apparent relaxation time graphs.

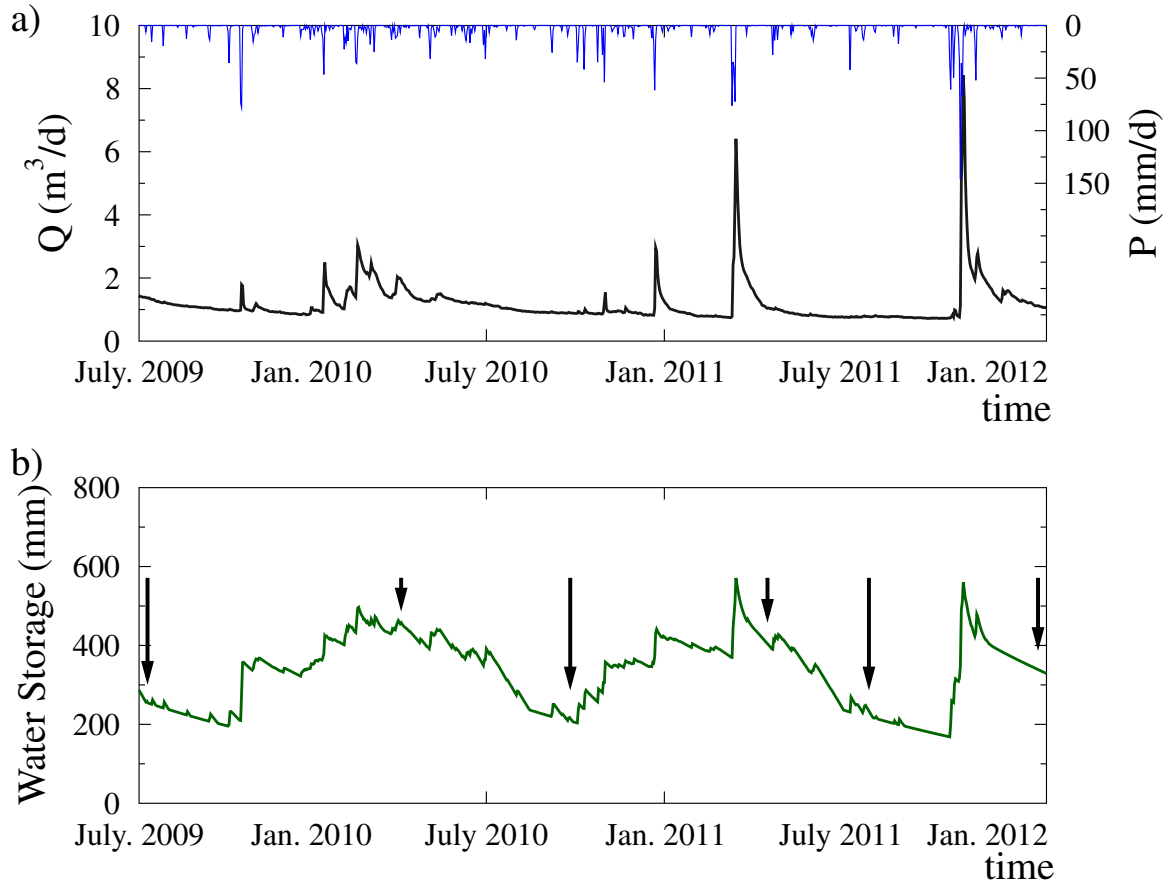


Figure 5: Water storage change on the Durzon catchment: a) Daily rainfall and measured Durzon spring discharge, b) Simulated water storage change on the Durzon catchment. Arrows stand for MRS field campaigns. The water storage at July 1st 2009 is arbitrarily set to zero.

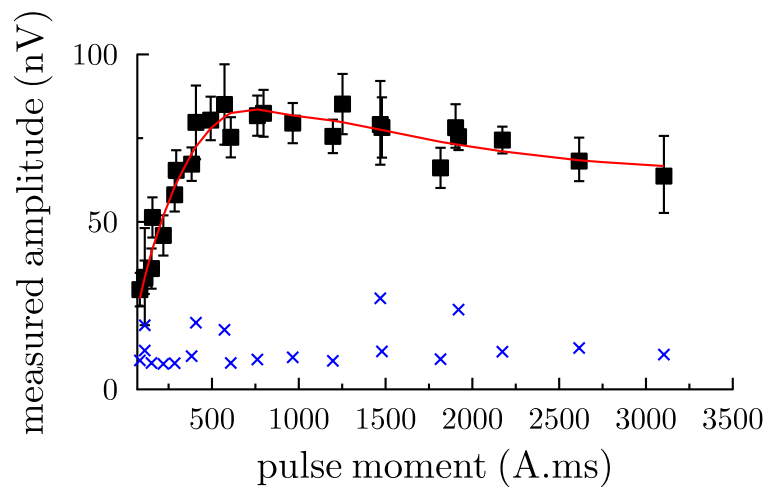


Figure 6: MRS signal and inversion results on H3 site. Square: signal amplitude e . Cross: ambient noise measurements. Line: MRS inversion.

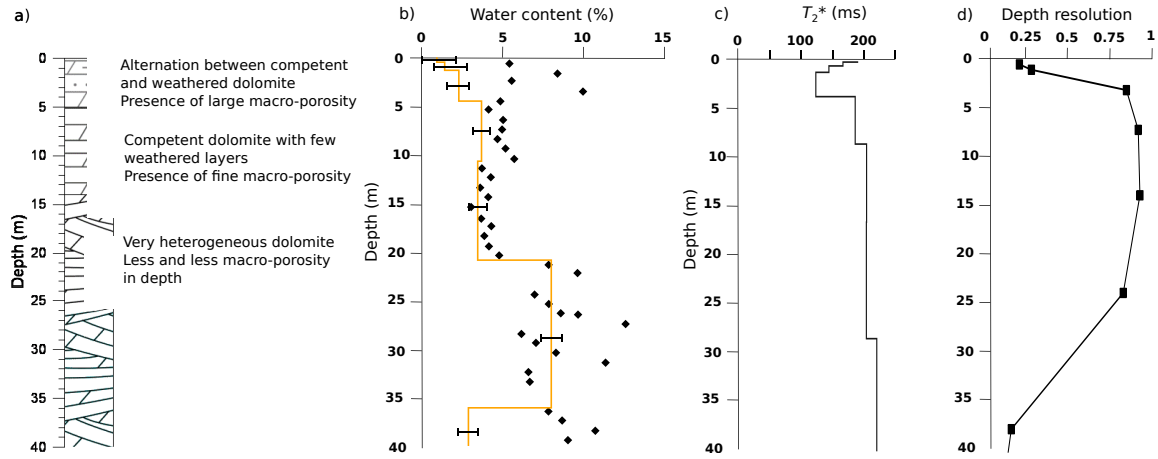


Figure 7: Comparison of MRS results and core samples from a drillhole performed at the center of H3 site: a) simplified log resulting from a synthesis of visual observations performed on 3 different boreholes, b) core sample water content (markers), and MRS (%) - line, c) $T_2^*(z)$ (ms), d) depth resolution of the MRS model.

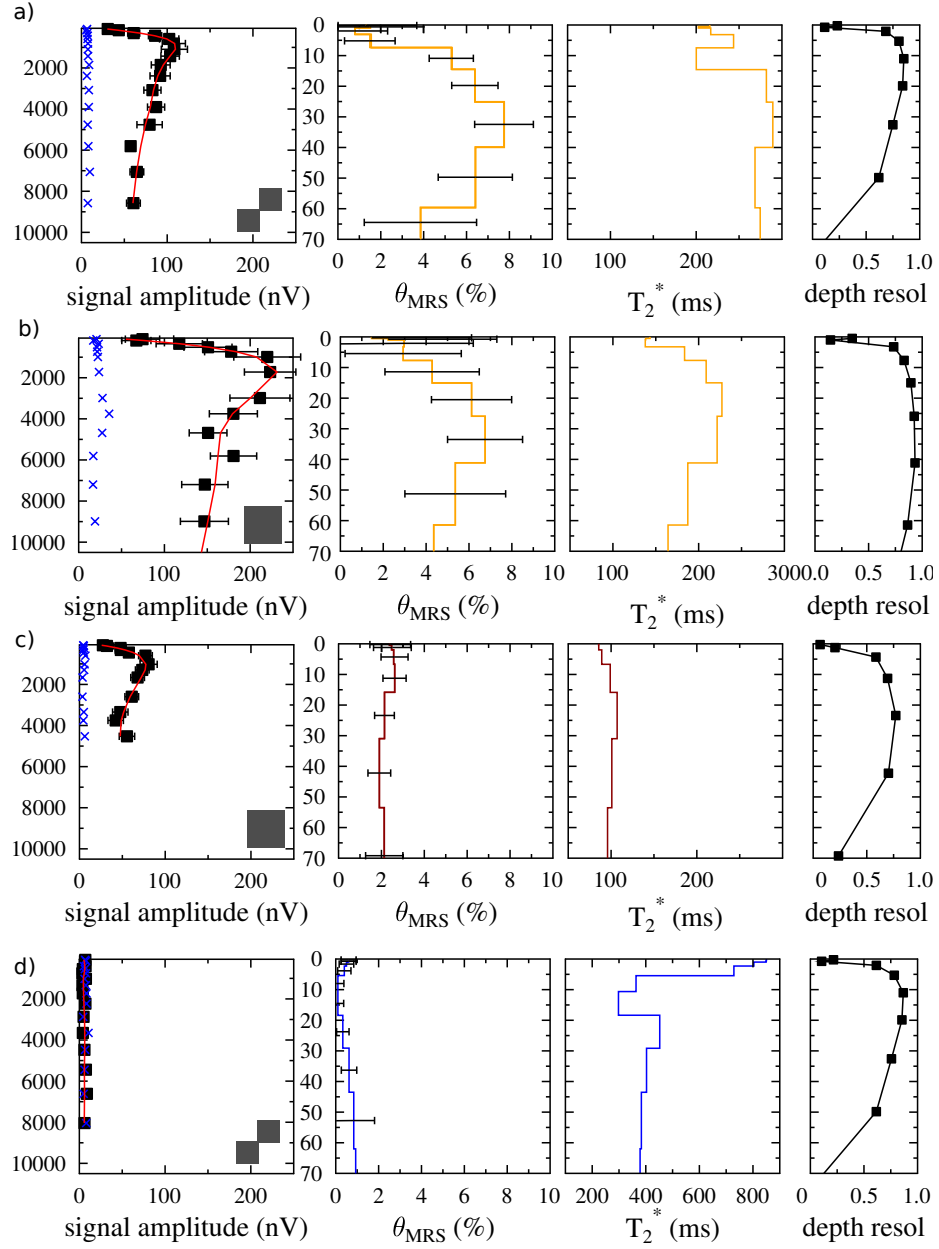


Figure 8: Example of MRS signal and inversion for selected sites: a) Chou (ruiniform dolomite, eight square loop), b) Cana (ruiniform dolomite, square loop), c) SambH (siliceous limestone, square loop), d) CombR (marly limestone, eight square loop). MRS results plots (left): signal amplitude ϵ (squares), ambient noise measurements (crosses), MRS inversion (line).

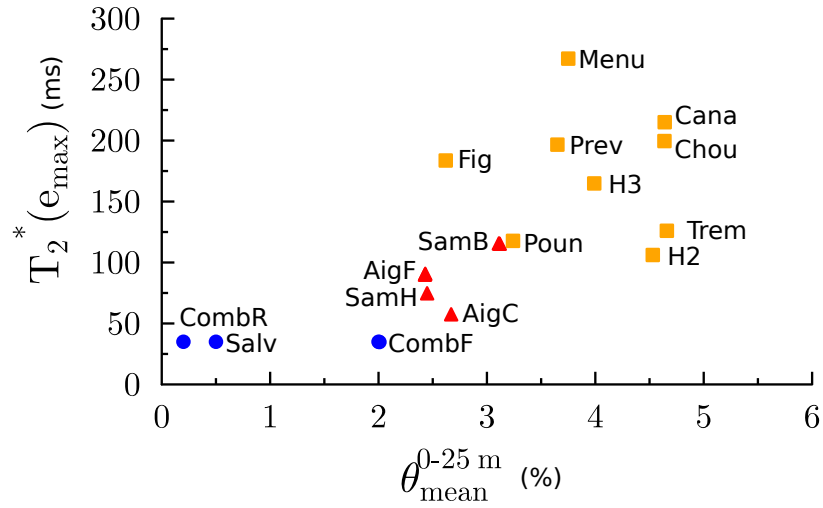


Figure 9: $T_2^*(e_{max})$ as a function of θ_{mean}^{0-25m} for soundings performed in ruiniform dolomite (squares), siliceous limestones (triangles), and marly limestone (circles). For marly limestones, $T_2^*(e_{max})$ is set to 35 ms referring to the instrumental dead time of the device, and θ_{mean}^{0-25m} is set equal to the maximum water content.

Site number	Site name	Lith.	Fract.	Karst.	Slope	Clay	V.	Temp. lake
1	AigC	SL	Y	N	> 5%	Y	C	Far
2	AigF	SL	N	N	> 5%	Y	POW	Far
3	Cana	RD	N	N	< 5%	N	PW	Far
4	Chou	RD	N	N	< 5%	N	EG	Close
5	CombF	RD	N	N	> 5%	N	EG	Far
6	CombR	ML	N	N	> 5%	N	OG	Far
7	Fig	RD	N	N	< 5%	Y	C	Far
8	H2	RD	N	Y	< 5%	Y	C	Far
9	H3	RD	N	N	< 5%	N	OG	Far
10	Menu	RD	N	N	< 5%	Y	OG	Close
11	Poun	RD	Y	N	< 5%	N	EG	Far
12	Prev	RD	N	N	< 5%	N	C	Within
13	Salv	ML	N	N	> 5%	N	OG	Far
14	SamB	SL	N	N	< 5%	Y	C	Far
15	SamH	SL	N	N	> 5%	Y	C	Far
16	Trem	RD	N	Y	< 5%	N	EG	Far

Table 1: Synthesis of the sites characteristics. Lith.: lithology (ML: marly limestone. SL: siliceous limestone. RD: ruiniform dolomite). Fract. /Karst. : fractures / vertical karstification features (Y: presence. N: absence). Clay : clay cover (Y: presence. N: absence). V.: vegetation type (EG: encroached grassland, OG: open grassland, PW: pine wood, POW: pine and oak wood, C: cropland). Temp. lake: location relative to the temporary lake area.

Site number	Site name	Freq.	S/N	EN/IN	Max. depth resol. (m)	Field Camp.	Loop shape	Nb. of stacks
1	AigC	1973.2	3.3	1.4	43	12/01	Eight	200
2	AigF	1972.0	2.2	1.6	55	12/01	Square	300
3	Cana	1973.5	4.6	4.6	59	12/01	Square	120
4	Chou	1970.1	8.3	1.6	45	12/01	Eight	120
5	CombF	1970.4	1.4	3	-	11/08	Eight	170
6	CombR	1973.2	1.0	1.4	-	11/08	Eight	130
7	Fig	1970.1	2.5	2.6	38	11/08	Eight	300
8	H2	1971.6	2.1	3.6	42	09/07	Eight	200
8	H2	1971.6	3.5	2.1	42	10/04	Eight	250
8	H2	1971.6	8.9	1.2	42	10/10	Eight	110
8	H2	1971.6	3.9	2.0	42	11/05	Eight	100
9	H3	1971.6	6	1.5	32	09/07	Eight	150
9	H3	1971.6	3	3.2	32	10/04	Eight	150
9	H3	1971.6	5.2	1.9	32	10/10	Eight	100
9	H3	1973.2	3.7	4.2	32	11/05	Eight	100
9	H3	1973.2	3.2	2.9	32	12/01	Eight	150
10	Menu	1970.1	2.4	3.2	38	09/07	Eight	200
10	Menu	1971.6	7	2.6	38	10/04	Eight	250
10	Menu	1971.6	8	1.4	38	10/10	Eight	160
10	Menu	1972.0	2.5	3.6	38	11/05	Eight	180
11	Poun	1970.1	5.6	1.3	42	12/01	Eight	120
12	Prev	1973.2	12.8	1.8	42	12/01	Square	200

13	Salv	1973.2	1.0	2.2	-	11/08	Eight	200
14	SamB	1973.2	3.2	2.1	53	12/01	Square	180
15	SamH	1973.2	4.9	1.0	25	12/01	Square	180
16	Trem	1973.2	8.3	1.5	45	12/01	Eight	120

Table 2: Synthesis of the MRS investigations. Freq.: signal frequency. S/N: signal-to-noise ratio. EN/IN: external noise to instrumental noise ratio. Max. depth resol.: Maximum resolution depth

Site	Site	r	$\theta_{\text{MRS}}^{0-10\text{m}}$	$\theta_{\text{MRS}}^{10-40\text{m}}$	$T_2^{*0-10\text{m}}$	$T_2^{*10-40\text{m}}$
No.	name	(nV)	(%)	(%)	(ms)	(ms)
1	AigC	4.6	2.05	5.09	51	73.17
2	AigF	3.8	2.09	3.18	34	76
3	Cana	11.1	3.97	5.96	163	214
4	Chou	5.2	1.51	7.55	595	221
5	CombF	-	-	-	-	-
6	CombR	-	-	-	-	-
7	Fig	4.2	2.14	4.25	22	174
8	H2	4.9	3.17	6.12	19	126
9	H3	5.9	2.88	4.65	145	173
10	Menu	4	2.73	3.84	116	374
11	Poun	3.4	2.16	5.07	176	112
12	Prev	6.5	5.17	3.93	133	171
13	Salv	-	-	-		
14	SamB	4.4	3.17	6	25	148
15	SamH	4.9	3.55	1.73	70	149
16	Trem	4.4	3.42	5.52	143	512

Table 3: Synthesis of the MRS inversions. r : model residuals. $\theta_{\text{MRS}}^{0-10\text{m}}$, $\theta_{\text{MRS}}^{10-40\text{m}}$, : average MRS water content for the 0-10 and 10-40 m respectively. $T_2^{*0-10\text{m}}$, $T_2^{*10-40\text{m}}$: average T_2^* for the 0-10 and 10-40 m respectively.

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